

# 1 **First dinosaur remains from Ireland**

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9

## 10 **ABSTRACT**

11 Several specimens from the Lias Group (Lower Jurassic) of Northern Ireland have been  
12 suspected as dinosaurian in origin. Bone histology and morphology demonstrates that two  
13 of these, both from the same locality in Co. Antrim, demonstrably are from dinosaurs.  
14 We interpret one as the proximal end of the left femur of a basal thyreophoran  
15 ornithischian, and tentatively assign it to cf. *Scelidosaurus*. The other is the proximal part  
16 of the left tibia of an indeterminate neotheropod, perhaps a stem-averostran similar to  
17 *Sarcosaurus*, or a megalosauroid. These are the first dinosaur remains reported from  
18 anywhere in Ireland and some of the most westerly in Europe, and are among only a  
19 small number of dinosaurs known from the Hettangian Stage. Two additional specimens  
20 are no longer considered to be from dinosaurs. We interpret one as a surangular or  
21 mandible fragment from a large marine reptile, perhaps an ichthyosaur or pliosaur; the  
22 other is a polygonal fragment of Paleocene basalt.

23

24 **Key words:** Ornithischia, Theropoda, Early Jurassic, Hettangian, Northern Ireland.

25

## 26 **1. Introduction**

27 Although dramatic finds from North America, Asia and elsewhere have dominated the

28 literature for decades, Britain remains one of the richest areas in the world in terms of  
29 documented dinosaur diversity. Important faunas have been recovered from the Upper  
30 Triassic (e.g. Benton et al. 2000, Whiteside et al. 2016), the Jurassic of south central and  
31 eastern England (Benton and Spencer 1995) and especially the Lower Cretaceous of the  
32 Isle of Wight and the Weald (Martill and Naish 2001), but elsewhere in the UK dinosaur  
33 remains are less abundant. Footprints and skeletal remains have been discovered in the  
34 Jurassic of Yorkshire (Whyte et al. 2010), in northern England; from the Late Triassic  
35 and Early Jurassic of south Wales (Benton and Spencer 1995, Yates 2003, Martill et al.  
36 2016); and the Jurassic of both western and eastern Scotland (Andrews and Hudson 1984,  
37 Benton et al. 1995, Clark 2018, dePolo et al. 2018, Young et al. 2018). However, no  
38 dinosaur remains have hitherto been recorded from the island of Ireland.

39 Some might attribute the apparent absence of dinosaur remains from Ireland to the  
40 activities of St Patrick, whose apparent success in casting out snakes is well known, but  
41 there is a more mundane explanation. The sparsity of dinosaur remains in Britain beyond  
42 southern England directly reflects the geology of these regions. Suitable Mesozoic rocks  
43 are poorly represented in northern and western Britain where most rocks are too old to  
44 contain dinosaur remains, and this is true too across large areas of Ireland. Outside of  
45 Northern Ireland there are just a few, largely concealed, patches of Mesozoic including  
46 Triassic clastics in the Kingscourt Inlier (Simms 2009) and possibly in the Wexford  
47 Outlier (Clayton et al. 1986), Jurassic clastics near Cork (Higgs and Beese 1986) and  
48 Carrick-on-Suir (Higgs and Jones 2000), and a tiny patch of Upper Cretaceous Chalk in  
49 Co. Kerry (Walsh 1966).

50 Substantial Mesozoic basins are present in the north-east (Larne-Lough Neagh Basin) and  
51 north-west (Foyle-Rathlin Basin) of Northern Ireland but they are mostly concealed  
52 beneath younger rocks and exposure is confined largely to a narrow strip around the  
53 margins of the Paleocene basalt plateau (Fig. 1). Triassic, Jurassic and Cretaceous strata  
54 are all represented (Fig. 2) but there is a gap, representing at least 100 million years,  
55 between the youngest preserved Jurassic strata (early Pliensbachian) and oldest preserved  
56 Cretaceous (early Cenomanian), caused by pre-Cenomanian erosion of earlier Mesozoic  
57 strata, and in places the Jurassic is cut out altogether where Cretaceous strata rest directly  
58 upon Triassic or even older rocks.

59 However, despite being of the right age to contain dinosaurs these Mesozoic strata are  
60 developed in facies that do not favour the preservation of terrestrial vertebrates. The  
61 Jurassic and Cretaceous strata here are entirely of open marine facies in which dinosaur  
62 fossils are unlikely to be common, while the Upper Triassic succession is largely non-  
63 marine and developed in red-bed facies of the Mercia Mudstone Group (Milroy et al.  
64 2019) in which fossils of any kind are rarely preserved. In short, Ireland's rocks are either  
65 of the wrong age or of the wrong type to contain dinosaur fossils.

66 Palaeogeographic reconstructions (Cope et al. 1992) suggest that at times during the  
67 Mesozoic dinosaurs could have walked on dry land across parts of what is now Ireland.  
68 However, if any rocks were deposited here at these times they were either removed long  
69 ago by erosion, along with any dinosaur remains that they might contain, or they were  
70 deposited in environments distant from dinosaur habitats. Finding an Irish dinosaur might  
71 seem a hopeless task but, nonetheless, several potential candidates have been identified  
72 and are described for the first time here.

73

## 74 **2. Occurrence and history of discovery**

75 Several specimens from Northern Ireland, now held in the collections of National  
76 Museums NI, have been identified as potentially of dinosaurian origin. One of these  
77 (BELUM K1642) was acquired in 1920 from the Belfast Natural History and  
78 Philosophical Society, in existence since 1821. Information with the specimen states only  
79 that it is from the "Lower Lias, Glenarm, Co. Antrim" and that possibly it was collected  
80 in the late 19<sup>th</sup> or early 20<sup>th</sup> century by William Gray (1830-1917), a well-known  
81 collector of fossils and antiquities and an ardent campaigner for a municipal museum in  
82 Belfast. The Lias Group is not exposed in situ in the immediate vicinity of the town of  
83 Glenarm but it is seen in landslips adjacent to the coast road at Straidkilly, to the north-  
84 west of the town (NI grid ref. D305163; 54.976°N, -5.96°W), and Minnis North, to the  
85 south-east (NI grid ref. D337137; 54.951°N, -5.91°W), and this bone probably originates  
86 from one or other of these sites. At Straidkilly the succession extends no higher than the  
87 Planorbis Zone but at Minnis North it encompasses the entire Hettangian Stage, and  
88 perhaps the base of the succeeding Sinemurian Stage. However, without further

89 information we can state only that the source of this specimen is 'probably Hettangian'.  
90 The three remaining specimens (BELUM K3998, K12493, K2015.1.54) were recovered  
91 from among boulders and shingle on the beach near the Gobbins, on the east coast of  
92 Islandmagee, Co. Antrim (Grid ref. J485968; 54.799°N, -5.690°W), by the late Roger  
93 Byrne between 1980 and 2000. Along the Gobbins coast the Antrim Lava Group basalts  
94 (Paleocene) and Ulster White Limestone Formation (Cretaceous, Santonian to  
95 Campanian) have slumped on the less mechanically competent lithologies of the Lias  
96 Group (Jurassic, Hettangian) and Hibernian Greensands Formation (Cretaceous,  
97 Cenomanian to Coniacian) beneath. The succession is highly disturbed but small  
98 exposures at the rear of the beach and among the boulders on the foreshore expose sand  
99 and silt facies of the Hibernian Greensand Formation lying unconformably upon grey  
100 clays and subordinate muddy limestones of the Lias Group (Planorbis Zone) and the  
101 upper Cotham Member of the Penarth Group (Triassic, Rhaetian Stage).

102 The Gobbins beach itself, extending south from the famous Gobbins cliff path, is covered  
103 with rounded fragments of basalt and white limestone. Fossil material is sparse and often  
104 heavily abraded. The specimens found by Roger Byrne were ex situ and hence might  
105 originate from any level in the local Mesozoic succession, but comparison with other  
106 vertebrate material from the Northern Ireland Mesozoic is instructive in establishing their  
107 probable source. The prevailing colour of the bone in these specimens is a very dark  
108 brownish-black, similar to marine vertebrate material recovered from Lias Group  
109 mudstones at various sites in Northern Ireland but distinct from the dark chocolate brown  
110 colour of bones and teeth found in the Hibernian Greensand or the pale brown colour of  
111 those from the Ulster White Limestone. Furthermore, a small patch of grey indurated  
112 calcareous mudstone on one of the specimens (K3998), and a pale grey micrite infill of  
113 the medullary cavity in another (K12493), resemble Lias Group lithologies but are  
114 distinct from lithologies in the Penarth Group or Chalk Group (Ulster White Limestone  
115 and Hibernian Greensand formations) exposed hereabouts. Together these lines of  
116 evidence suggest that the Gobbins specimens are from the Lias Group, and probably from  
117 the Planorbis Zone at the base of the Hettangian Stage, or from Lias Group strata  
118 immediately beneath often referred to as the 'Pre-planorbis Beds'.

119 For many years the base of the Jurassic System in Britain was placed at the first  
120 appearance of the ammonite *Psiloceras planorbis*, or its assumed correlatives  
121 (Warrington et al. 1994, 2008) but in 2007 formal proposals were made to designate a  
122 GSSP for this boundary. The Triassic-Jurassic boundary section at Larne, Co. Antrim,  
123 was a candidate site (Simms and Jeram 2007) although it was the section at Kuhjoch, in  
124 Austria, that ultimately was selected (von Hillebrandt et al. 2013) based on the first  
125 occurrence of the ammonite *Psiloceras spelae tirolicum*. Direct biostratigraphic  
126 correlation between Northern Ireland and the Kuhjoch section is not possible but the  
127 proposed Stage and System boundary, as defined in Austria, corresponds to a significant  
128 negative excursion in the carbon isotope curve (Ruhl et al. 2009) that has been recognised  
129 in the Larne succession. It indicates that the base of the Jurassic System lies within the  
130 ‘pre-Planorbis beds’ of the Lias Group here, in the lower part of bed 22 (Simms and  
131 Jeram 2007) and about three metres below the first appearance of ammonites (*Psiloceras*  
132 *erugatum*) in bed 24, although recent work by Hodges (in press) suggests that the Jurassic  
133 may commence lower in the succession.

134

### 135 **3. Early Jurassic dinosaurs in western Europe**

136 Dinosaur remains are uncommon in the Lower Jurassic of western Europe so even the  
137 fragmentary remains described here are potentially of significance. The best known is the  
138 thyreophoran ornithischian *Scelidosaurus harrisoni* from the Sinemurian of the Dorset  
139 coast, England. The material on which Owen (1861) based his original description  
140 actually encompassed two distinct animals; an armoured ornithischian represented by a  
141 near complete skeleton, and a theropod represented by part of a hind limb. Bizarrely it  
142 was the latter that was originally designated as the holotype, although it is the armoured  
143 ornithischian that now holds the name. Several near complete skeletons of *Scelidosaurus*  
144 are now known, all from Dorset. A second questionable thyreophoran, *Lusitanosaurus*  
145 *liasicus* has been described from the Sinemurian of Portugal (Lapparent and Zbyszewski,  
146 1957) and another, *Emausaurus ernsti*, is from the Toarcian of northern Germany  
147 (Haubold 1990). Remains of several theropods are known from the Hettangian;  
148 *Lophostrophus airelensis* (Cuny and Galton 1993) from Airel, Normandy; *Sarcosaurus*

149 *woodi* from central England (Andrews 1921, Woodward 1908, von Huene 1932, Ezcurra  
150 et al. 2020); *Dracoraptor hanigani* from Lavernock, south Wales (Martill et al. 2016); a  
151 pedal phalanx and tooth fragment from Luxembourg (Delsate and Ezcurra 2014); and,  
152 from the Sinemurian, the partial hind limb to which the name *Scelidosaurus harrisoni*  
153 was first applied. A sauropodomorph hind limb, *Ohmdenosaurus liasicus*, has also been  
154 described from the Toarcian of Baden-Württemberg (Wild 1978, Stumpf et al. 2015).  
155 Much of this material is specifically indeterminate and even some of the named taxa,  
156 such as *Ohmdenosaurus*, are so fragmentary that they are perhaps best considered *nomina*  
157 *dubia*.

158 Slightly older taxa, from the late Triassic (Rhaetian), include a theropod dentary,  
159 ?*Megalosaurus cambrensis*, from south Wales (Newton 1899, Galton, 2005); an  
160 incomplete melanosaurid sauropodomorph skeleton, *Camelotia borealis*, from the  
161 Penarth Group of Somerset in south-west England (Galton 1985); various taxonomically  
162 indeterminate fragments from the Westbury Formation of Newark, Nottinghamshire  
163 (Martill and Dawn 1986) and Aust Cliff, near Bristol (Storrs 1994); and the basal  
164 sauropodomorphs *Thecodontosaurus* and *Pantydraco* from supposedly Rhaetian fissure  
165 fills in south-west England and south Wales (Galton et al. 2007, Riley and Stutchbury  
166 1836, Whiteside et al. 2016). Supposed dinosaur limb bone fragments from the Penarth  
167 Group at Aust Cliff, near Bristol (Galton 2005), are probably jaw fragments from giant  
168 ichthyosaurs (Lomax et al. 2018).

169

#### 170 **4. Evidence for dinosaurs in Ireland**

171 In Northern Ireland the Jurassic succession is entirely marine. Bones of marine fish and  
172 reptiles, ichthyosaurs and plesiosaurs, have been recovered from many exposures of the  
173 Lower Jurassic here and so our initial assumption must be that those bones we consider as  
174 potentially dinosaurian in origin are, in fact, from marine vertebrates. This possibility can  
175 largely be discounted. Firstly, no Lower Jurassic fish yet discovered attains a size  
176 comparable with that indicated by these bones. Secondly, the shape of each of the bone  
177 fragments in question does not immediately resemble any of the bones found in  
178 ichthyosaur or plesiosaur skeletons, although caution must be exercised here as

179 fragmentary ichthyosaur bones have in the past been mistaken for those of dinosaurs  
180 (Lomax et al. 2018). Finally, and most significantly, the fine structure of these bones  
181 differs significantly from that found in fish, ichthyosaurs or plesiosaurs/pliosaur.

182

#### 183 **4.1. *Bone histology and surface texture***

184 The bones of Jurassic ichthyosaurs and plesiosaurs, including their long bones and  
185 vertebrae, are commonly dominated by a porous or cancellous structure, with little or no  
186 medullary cavity, surrounded by a compact peripheral layer (Houssaye et al. 2014, Liebe  
187 and Hurum 2012) that typically is relatively thin (less than 20% of bone radius). This is  
188 similar to extant cetaceans and reflects their highly evolved aquatic mode of life where  
189 overall density of the animal is more critical than the load-bearing properties of the  
190 bones. In contrast the load-bearing bones of terrestrial vertebrates, such as dinosaurs, tend  
191 to have a more thickened (30-40% of bone radius) outer zone of compact bone around an  
192 inner core of cancellous bone and/or a medullary cavity which, in life, contains bone  
193 marrow (e.g. Hübner 2012, Redelstorff et al. 2013). A further significant distinguishing  
194 character is that the surface texture of dinosaur long bones is very smooth whereas those  
195 of marine reptiles, including long bones, ribs and jaws, commonly tend to show a greater  
196 rugosity in the form of longitudinal striations subparallel to the bone long axis.

197 The typically dinosaurian aspects of bone histology seen in K3998 were recognised as  
198 early as 1989 by the late Robin Reid of Queens University Belfast, who had developed an  
199 interest in dinosaur bone histology (e.g. Reid 1984), and the discovery of this specimen  
200 was reported in the local press. It was Roger Byrne himself that recognised the second  
201 dinosaur bone fragment from the Gobbins (K12493) based essentially on the same  
202 criteria (dense outer bone and large medullary cavity). Two other potential candidates  
203 (K1642 and K2015.1.54) have in the past been suggested, or suspected, as being of  
204 dinosaurian origin but a closer reexamination reveals that this is not so.

205

#### 206 **4.2. *Morphological descriptions***

207 Bone histology and surface texture can be used to establish if these fragments are from

208 dinosaurs, but it is their shape that we have used in our attempt to identify which bones  
209 they are and from broadly which group(s) of dinosaurs they originate. Robin Reid  
210 suggested in 1989 (Ulster Museum, unpublished notes) that K3998 was the proximal end  
211 of a femur, perhaps from the well-known early Jurassic thyreophoran genus  
212 *Scelidosaurus* although his reason for assigning it to that genus may have been based on  
213 little more than that it being the best-known early Jurassic dinosaur. With the second  
214 fragment (K12493) found at the same site it was, understandably, assumed to be from the  
215 same animal as K3998. However, on the basis of their size difference alone it is  
216 immediately apparent that this is not the case and the two bone fragments clearly  
217 represent different individuals and distinct taxa.

218

### 219 *4.3. Specimen descriptions and interpretation*

#### 220 *4.3.1. Specimen BELUM K3998a, b*

221 This specimen (Figure 3a-j) was collected by Roger Byrne on 13<sup>th</sup> January 1980 and was  
222 the first Irish dinosaur bone to be recognised, by Dr Robin Reid in 1989. It is a short  
223 fragment of a bone, smooth on its outer surface, abraded at its presumed proximal end  
224 and sectioned distally to reveal a thickened bone wall sharply demarcated from a lumen  
225 with a very open trabecular bone fill (Fig. 3c, f). There is no evidence of a medullary  
226 cavity, perhaps reflecting the proximity of this cut surface to the original bone  
227 termination. A glass-mounted thin section of this cut surface is in the Ulster museum  
228 collections (BELUM K3998b). The original preserved extent of the bone beyond the cut  
229 surface is not known, but it is not believed to be substantial. The outline of the bone, as  
230 seen laterally, is roughly quadrate. One face is broadly concave with a prominent sulcus  
231 offset to one side and parallel to the long axis of the bone (Fig. 3b, e). On the other side  
232 there is a prominent boss, located slightly off-centre, that is higher towards the middle  
233 and fades away toward two of the margins of the subquadrate outline (Fig. 3a, d).

234 Surface texture and histology establish this specimen as a dinosaur bone and we interpret  
235 it as the proximal end of a left femur, sectioned across its distal end and with the  
236 proximal end abraded and missing the femoral head. What is preserved of the femoral  
237 shaft is almost straight, but there is too little remaining to say anything more about shaft



238 morphology. The heavily worn base of the anterior trochanter is present on the  
239 anterolateral surface of the femoral shaft (Fig. 3a, d, h). A small area of matrix, on which  
240 the accession number is written, remains between the proximal preserved part of the  
241 anterior trochanter and the femoral shaft (Fig. 3g) and demonstrates that they were well  
242 separated from each other. There is no evidence to suggest that the femoral head was  
243 separated from the greater and anterior trochanters by a constriction. The anterior  
244 trochanter joins with the femoral shaft via a distally expanding ridge that extends along  
245 the preserved portion of the anterolateral edge of the element. A finger-like mound  
246 extends proximodistally along the anteromedial surface of the shaft (Fig. 3b, e, h), and  
247 presumably this structure was confluent or merged with the distal margin of the femoral  
248 head. It suggests that the femoral head would have emerged gently from the shaft of the  
249 femur.

250 The presence of a prominent anterior trochanter, well separated from the femoral shaft, is  
251 further evidence that this specimen is from a dinosaur, while thickening of the bone wall  
252 with trabecular bone is characteristic of ornithischians. Furthermore, the lack of any  
253 indication of a constriction separating the femoral head from the greater and anterior  
254 trochanters is a plesiomorphic trait shared by basal ornithischians to the exclusion of  
255 more highly derived forms (Butler et al. 2012). BELUM K3998 is relatively large for an  
256 Early Jurassic ornithischian, with ornithischian diversity at this time characterised by  
257 small-bodied heterodontosaurids (Sereno, 2012) and basal thyreophorans (Norman et al.  
258 2004), with no heterodontosaurid known that approaches a size even close to that of  
259 BELUM K3998 (Sereno 2012). The quadrate outline of the femoral shaft of BELUM  
260 K3998 in cross-section is also quite unlike the subcircular and elliptical femoral shaft  
261 cross-sections seen in other Early Jurassic ornithischians such as *Lesothosaurus* and  
262 heterodontosaurids (Butler 2010, Norman 2019), but it does resemble that of the basal  
263 thyreophoran *Scelidosaurus*. Among Early Jurassic ornithischians only *Scelidosaurus* and  
264 *Lusitanosaurus* are known to attain a size comparable with that indicated by BELUM  
265 K3998 (Antunes and Mateus 2003, Norman 2019), and comparison with femora of  
266 *Scelidosaurus* (NHMUK R1111 and NHMUK R6704) does show remarkable similarities  
267 (Figure 3h, k). The size and morphology of BELUM K3998, the positioning and  
268 condition of the anterior trochanter, and the finger-like ridge present on its anteromedial

269 surface, all match closely with *Scelidosaurus* and differ from what is seen in the  
270 significantly smaller basal thyreophoran *Scutellosaurus* (Norman 2019, Rosenbaum and  
271 Padian 2000).

272

#### 273 4.3.2. *Specimen BELUM K12493.*

274 This bone fragment (Fig. 4a-h, k, l) was collected by Roger Byrne on 15<sup>th</sup> April 1981. It  
275 is a portion of the shaft of a long bone with a smooth outer surface, abraded proximal  
276 articular surface and truncated distal section. It displays a robust lateral crista and has a  
277 triangular outline proximally with a sub-rectangular cross section distally (Fig. 4c, d).  
278 The proximal cross section is damaged, probably due to post-burial compaction, and  
279 reveals a thickened bone wall sharply demarcated from a lumen containing vacuous  
280 trabecular bone, with voids tending to be larger toward one side (Fig. 4c, k). The distal  
281 cross section reveals a somewhat thicker bone wall with a subcircular lumen (about 50%  
282 of bone diameter) that lacks trabecular bone. Most of this cavity is filled with pale grey  
283 micritic limestone but about the remaining third contains a pale, off-white, geopetal  
284 calcite infill containing small angular bone shards (Figure 4d, l). In both cross-sections  
285 the thickened bone displays a number of lines of arrested growth (LAGs) suggesting  
286 development under a seasonal climate (*sensu* Köhler et al. 2012). There is a flattened to  
287 weakly concave surface along one length of the diaphysis that we presume was occupied  
288 by an adjacent bone, as in the case of a tibia-fibula or metatarsal-metatarsal contact.

289 The surface texture and histology of this specimen establish it as a dinosaur bone, and we  
290 interpret it as the worn and truncated proximal portion of a left tibia. The shaft, as  
291 preserved, is straight although too little remains to infer much about its intact form. The  
292 distal portion of the cnemial crest rises from the anterior surface of the shaft (Fig. 4a, e)  
293 but neither medial nor lateral condyles are evident in this specimen, which we attribute to  
294 these structures being confined to the, now eroded, proximal-most region of the posterior  
295 side of the tibia. A large fibular crest extends proximodistally along the lateral surface of  
296 the tibial shaft (Fig. 4b, f). It is somewhat expanded anteroposteriorly at the midpoint of  
297 the preserved shaft and tapers towards the proximal and distal ends of the element. It was  
298 most likely lens or tear-drop shaped in lateral view but, despite significant wear, there is

299 no evidence of any internal bone texture exposed on the low broad crest suggesting that  
300 in life it did not rise to a sharp point.

301 BELUM K12493 is a heavily abraded bone fragment lacking terminations and hence  
302 identification is difficult. It does show some similarities with a proximal metatarsal 4 in  
303 the Early Cretaceous theropod *Neovenator salerii* Hutt, Martill and Barker 1996, from the  
304 Isle of Wight, but this is not an especially good match and it bears little similarity to the  
305 same elements in the Late Jurassic theropod *Allosaurus fragilis* Marsh 1877 (see Madsen  
306 1976 fig. 25A). Furthermore, were this the case the size of BELUM K12493 would imply  
307 the existence of an Early Jurassic theropod substantially larger than any currently known.  
308 Instead its morphology is more consistent with that of a tibia, with the combination of its  
309 size, prominent fibular crest, thin bone wall distally with an absence of trabecular bone,  
310 and the probable basis for a large cnemial crest, attesting to the theropodan affinities of  
311 this specimen. The fibular crest in particular is diagnostic, extending well into the  
312 proximal end of the tibia on BELUM K12493. This is the basal condition for theropods  
313 and is present in coelophysoids (Spielmann et al. 2007), dilophosaurids (Welles 1984),  
314 stem-averostrans such as *Sarcosaurus* (Ezcurra et al. 2020), as well as ceratosaurians  
315 (Madsen and Welles 2000, O'Connor 2007). Among tetanurans only some  
316 megalosauroids are known to possess such a proximally extended fibular crest (Benson  
317 2010a, Madsen 1976, Malafaia et al. 2018). Of these megalosauroids only a few develop  
318 a crest with a low anteroposteriorly broad mound, as seen in BELUM K12493, including  
319 *Megalosaurus*, *Piatnitzkysaurus*, *Condorraptor*, and *Magnosaurus* (Benson 2010a,  
320 2010b, Rauhut 2005). In other megalosauroids, including *Eustreptospondylus*,  
321 *Spinosaurus*, *Suchomimus* and *Torvosaurus* the fibular crest forms a narrow ridge, more  
322 typical of tetanurans (Benson 2010a, Ibrahim et al. 2014, Sadleir et al. 2008). The  
323 metriacanthosaurid *Sinraptor* also possesses an oval shaped, mound-like fibular crest, but  
324 it is positioned more anteriorly on the tibial shaft and forms a more prominent ridge than  
325 in either megalosauroids or BELUM K12493 (Currie and Zhao 1993).

326 The presence of an anteroposteriorly broad, proximally extended fibular crest on BELUM  
327 K12493 is similar to those found on some megalosauroids, specifically the  
328 Megalosauridae or Piatnitzkysauridae. It resembles the tibiae of *Magnosaurus* and  
329 matches closely those of the Bathonian *Megalosaurus bucklandi* itself (particularly

330 NHMUK PV R 12557 and OUMNH J.13562), although the latter are substantially larger  
331 than that represented by BELUM K12493 (Figure 4i, j). However, BELUM K12493  
332 differs from the tibia of these megalosauroids in other ways besides overall size. While  
333 the fibular crest of megalosauroids emerges gently from the tibial shaft to form a low  
334 mound, the crest on BELUM K12493 emerges from the tibial shaft much more abruptly.  
335 Additionally, the distal termination of the fibular crest in BELUM K12493 is more abrupt  
336 than those of megalosauroids which grade more gently into the bone shaft.

337 In these features BELUM K12493 bears a closer resemblance to *Sarcosaurus woodi* from  
338 the Lias Group (Hettangian–Sinemurian) of central England (*sensu* Ezcurra et al. 2020).  
339 *Sarcosaurus* was originally referred to Megalosauridae (Andrews 1921, von Huene  
340 1932), but subsequent analyses of the taxon have variously placed it among  
341 Coelophysidae/Coelophysoidea (Welles 1984; Carrano et al. 2005), basal Ceratosauria  
342 (Ezcurra 2012), Dilophosauridae (Dal Sasso et al. 2018) and, most recently, as stem-  
343 Averostran (Ezcurra et al. 2020). The size and morphology of BELUM K12493 and  
344 *Sarcosaurus* are consistent. They are recovered from approximately coeval strata within  
345 the Lias Group and their localities are within relatively close proximity (~375-390 km).  
346 Most likely, BELUM K12493 represents an animal similar to *Sarcosaurus*, but the  
347 limitations of the material and absence of shared autapomorphies precludes any definite  
348 referral to this taxon. We consider BELUM K12493 to represent an indeterminate  
349 neotheropod, likely representing either a stem-averostran or a megalosauroid.

350

#### 351 4.4. Additional specimens previously interpreted as dinosaurian

352 Two further specimens in the collections of the Ulster Museum have, in the past, been  
353 suggested as of dinosaurian origin but closer examination reveals these claims to be  
354 unfounded.

355

##### 356 4.4.1. Specimen BELUM K1642

357 This specimen (Figure 5a, g-h), which has been in the museum's collection for at least a  
358 century, is a straight, roughly cylindrical section of bone approximately 16 cm long and

359 6.5 cm across. It lacks any terminations, being broken obliquely at one end and  
360 terminated by a cut section at the other. In cross section it is roughly D-shaped, with the  
361 curved part of the D represented by the intact rounded outer surface of the bone and the  
362 straight section corresponding to a fractured surface exposing an area of cancellous bone  
363 (Figure 5g). The sectioned end of the bone (Figure 5h) reveals a relatively dense, yet still  
364 porous, outer layer extending through about 280° and interrupted by a zone of coarsely  
365 porous trabecular bone. The outer layer surrounds a relatively small (<40% of diameter)  
366 central cavity into which traces of coarse trabecular bone penetrate.

367 The bone was loaned to Dr Robin Reid in January 1980 and returned, 6 months later,  
368 with a section removed from one end but without comment from Dr Reid. He had  
369 developed an interest in dinosaur bone histology at this time and so we assume that this  
370 section was removed in relation to that project, on suspicion that it might be a dinosaur  
371 bone, but he did not convey his thoughts to the Ulster Museum geologists at the time.  
372 Superficially this fragment does resemble a dinosaur long bone shaft in its broadly  
373 cylindrical shape, albeit with about a fifth of its circumference apparently broken away,  
374 and its appearance in section of a distinct central lumen surrounded by relatively dense  
375 bone. However, closer examination of the specimen reveals features that are not  
376 consistent with this interpretation. The outer surface of K1642 is not smooth, as in  
377 dinosaur long bones, but has the striated appearance (Figure 5a, g) characteristic of the  
378 ribs, jaws and limb elements of ichthyosaurs (cf. fig 4 of Lomax et al. 2018) and  
379 plesiosaurs/pliosaurs. Furthermore, in section the outer bone layer is substantially less  
380 dense than is seen in dinosaurs, with a more conspicuously visible pattern of canals and  
381 LAGs (Figure 5h, cf. fig. 4d), and coarsely porous trabecular bone that occupies and  
382 extends into that section of the circumference where the outer surface appears to have  
383 been broken away. This is not what we would expect to find in any dinosaur long bone  
384 but invites comparison with the distal end of an ichthyosaur surangular bone, as described  
385 and figured by Lomax et al. (2018). Hence, we interpret this fragment as probably a  
386 fragment of surangular from a large ichthyosaur, perhaps *Temnodontosaurus*, or possibly  
387 a mandible fragment from a pliosaur such as the rhomaleosaurid *Atychodracon*.

388

389 4.4.2. Specimen BELUM K2015.1.54

390 In the year 2000 Roger Byrne discovered a small flat pentagonal object (Figure 5b-f) with  
391 a bone-like texture amongst shingle on the Gobbins beach. By association with the  
392 presumed scelidosaur femur (K3998) it was assumed that this object might be a  
393 scelidosaur osteoderm, although it lacked the point or curved crest typical of osteoderms  
394 in *Scelidosaurus harrisoni* (Norman 2019). The identity of this object remained  
395 enigmatic for almost two decades. Dinosaur experts were consulted but could see no  
396 evidence for dinosaurian affinity, and turtle experts similarly were baffled. It was only  
397 through examination by a fresh pair of eyes (DMM) in 2019 that the mystery was finally  
398 solved. It is not a bone at all but merely a small pentagonal piece of basalt! Who would  
399 have thought, in the land of the Giant's Causeway, that such a remarkably regular  
400 fragment of basalt could turn up on an Antrim beach? The lesson to be learned here is  
401 that some of us (MJS) were perhaps too influenced by the remarkably regular shape of  
402 this object and by its discovery on the same beach from which a dinosaur bone,  
403 tentatively assigned to *Scelidosaurus*, had been recovered.

404

## 405 5. Conclusions

406 Several specimens from Northern Ireland have been suspected, or suggested, as dinosaur  
407 bones but just two can be definitely assigned to this group on the basis of their bone  
408 histology, surface texture and morphology. Both are from the base of the fully marine  
409 Lias Group (Jurassic System, Hettangian Stage, probably Planorbis Zone) of the Gobbins,  
410 Islandmagee, Co. Antrim. Although fragmentary and beach-worn these bones retain  
411 sufficient characters for them to be identified and tentatively assigned to distinct clades.  
412 We interpret BELUM K3998 as the proximal end of a left femur of a thyreopophoran  
413 ornithischian, and tentatively assign it to cf. *Scelidosaurus*. If correct this underlines the  
414 close association of *Scelidosaurus* with the marine Lias Group and perhaps implies  
415 something about its environmental preference. We identify BELUM K12493 as the  
416 proximal portion of a left tibia from a neotheropod, and most likely from either a stem-  
417 averostran similar to *Sarcosaurus*, or from an indeterminate megalosauroid.

418 Reexamination of two other specimens reveals that neither are from dinosaurs. A bone

419 fragment from the Lias Group of Glenarm (BELUM K1642), suspected as being part of  
420 a dinosaur long bone, is probably a surangular or mandible fragment from a large marine  
421 reptile while an enigmatic pentagonal object (BELUM K2015.1.54) from the Gobbins,  
422 once thought to be a scelidosaur osteoderm, is nothing more than a small polygonal flake  
423 of Paleocene basalt!

424 Hettangian dinosaurs are rare and represented only by fragmentary material from England  
425 (Andrews 1921), Scotland (Benton et al. 1995) and Arizona (Lucas and Heckert 2001),  
426 and perhaps also from Morocco (Nicholl et al. 2018), South Africa (McPhee et al. 2015)  
427 and Antarctica (Smith et al. 2007). This global rarity emphasises the potential  
428 significance of even these fragmentary remains from Northern Ireland for understanding  
429 the evolution and biogeography of Early Jurassic dinosaurs.

430

#### 431 **Declaration of Competing Interest**

432 The authors declare that they have no known competing financial or personal  
433 relationships that could have appeared to influence the work reported in this paper.

434

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439 3D scans of the two Gobbins bones for the Portsmouth team.

440

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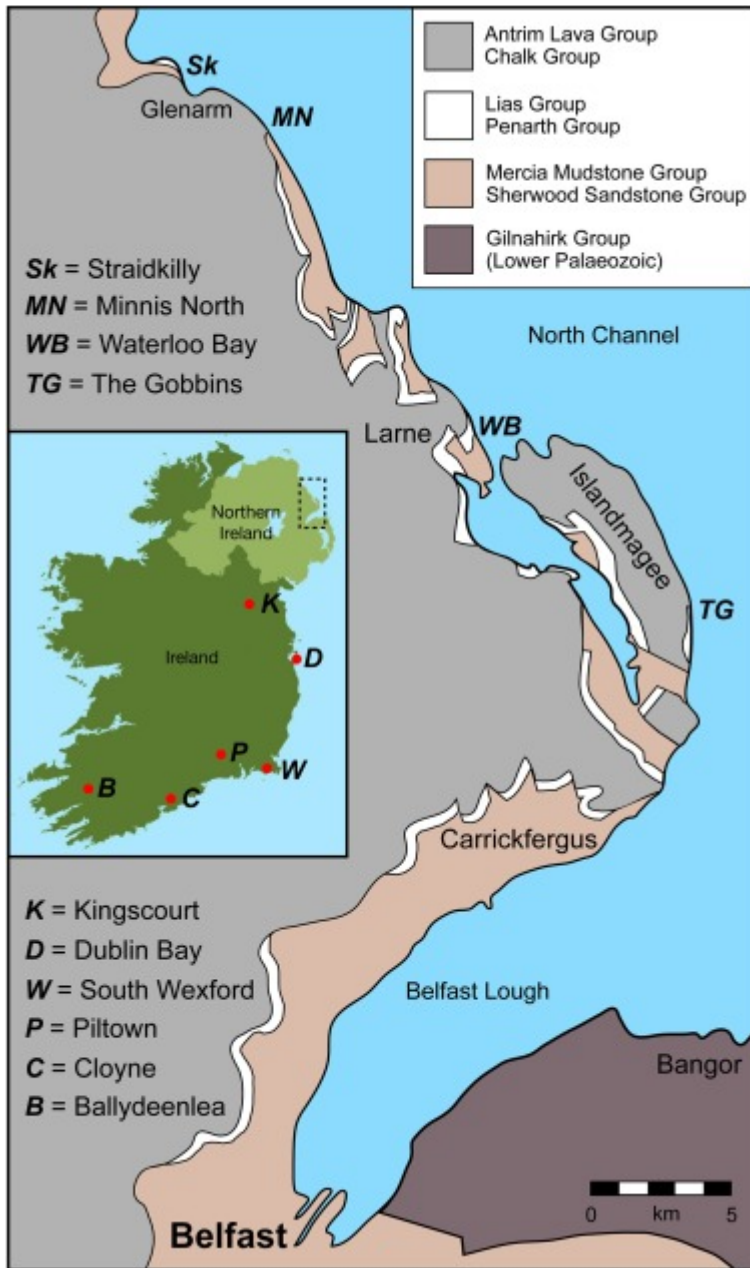
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659

660 **Figures**



661

662 **Fig. 1.** Map of eastern Co. Antrim, with generalised geology, for locations mentioned in  
 663 the text. MN = Minnis North; Sk = Straidkilly; TG = The Gobbins; WB = Waterloo Bay.

664 Inset: map of Northern Ireland showing location of area covered by geology map.

665



<b>CRETACEOUS</b>	<b>Chalk Group</b>	Maastrichtian to Santonian	Ulster White Limestone Formation	
		Santonian to Cenomanian	Hibernian Greensands Formation	
<b>JURASSIC</b>	<i>unconformity</i>	Sinemurian	Waterloo Mudstone Formation	Bucklandi Zone
				Angulata Zone
		Hettangian		Liasicus Zone
				Planorbis Zone
				'Pre-planorbis beds'
<b>TRIASSIC</b>	<b>Penarth Group</b>	Rhaetian	Lilstock Formation	
			Westbury Formation	
	<b>Mercia Mudstone Group</b>	Norian to Ladinian	Colin Glen Formation	
			undivided	
	<b>Sherwood Sandstone Group</b>	Ladinian to Scythian	undivided	

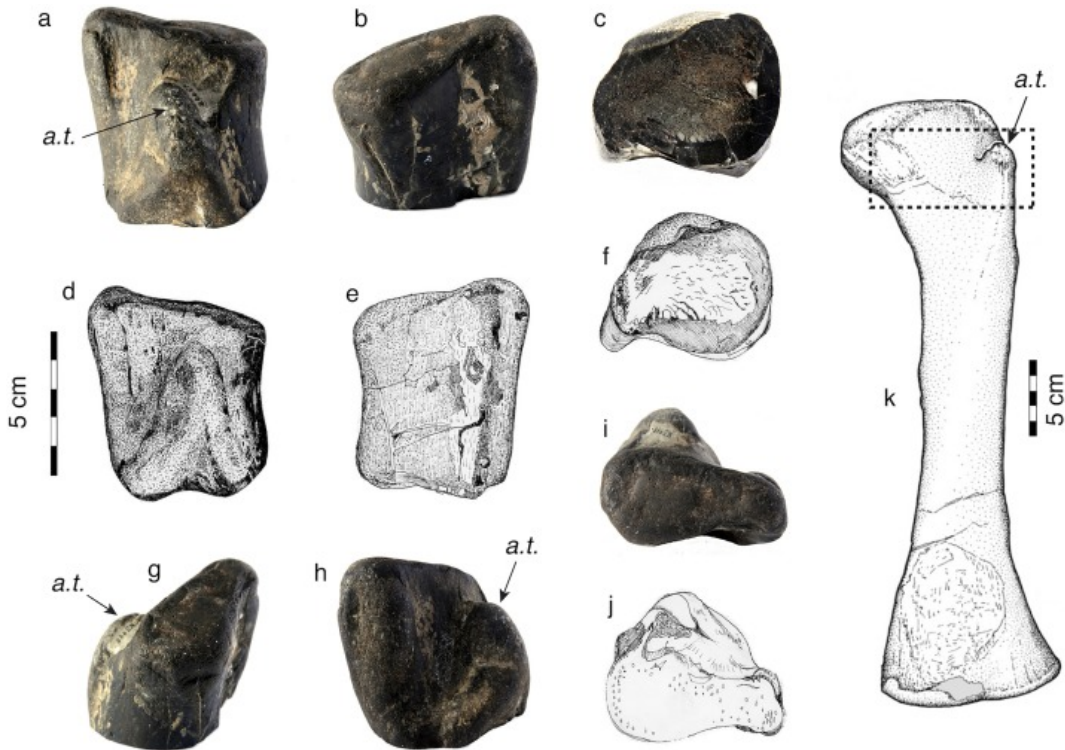
666

667 **Fig. 2.** Generalised litho- and chronostratigraphy for the succession exposed on the east  
668 Antrim coast.

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672

673 **Fig. 3.** Proximal fragment of left femur of thyreophoran ornithischian, cf. *Scelidosaurus*,  
674 BELUM K3998, from The Gobbins, Co. Antrim . Figures d-f, j drawn by Roger Byrne.  
675 a.t. = anterior trochanter.

676 **a, d.** Anterolateral view showing worn but conspicuous anterior trochanter at centre.

677 **b, e.** Posterior view

678 **c, f.** Sectioned surface of distal end with dense outer bones surrounding coarse trabecular  
679 bone.

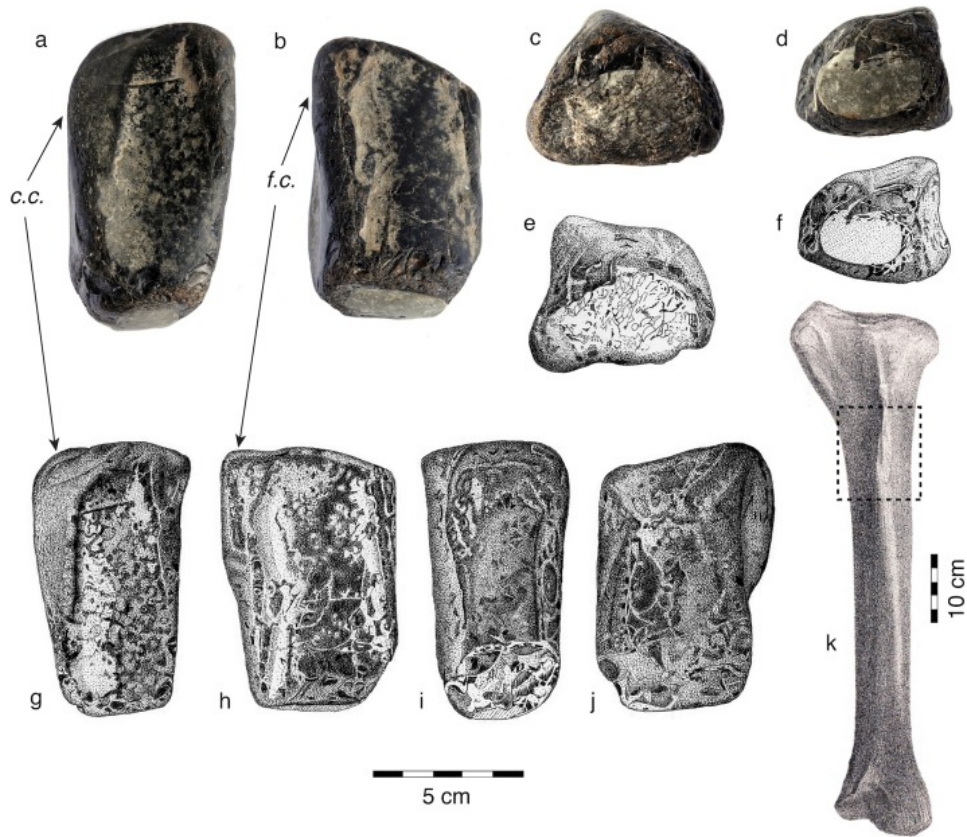
680 **g.** Lateral view, with anterior trochanter on left hand side, below accession number.

681 **h.** Medial view, with anterior trochanter conspicuous on right hand side.

682 **i, j.** Worn proximal end viewed from above.

683 **k.** Anterior view of left femur of *Scelidosaurus harrisoni* lectotype (NHMUK R1111),  
684 with section represented by K3998 outlined (from Norman 2019, fig. 78).

685



686

687 **Fig. 4.** Proximal fragment of left tibia of indeterminate neotheropod or megalosauroid  
 688 theropod, BELUM K12493, from The Gobbins, Co. Antrim. Figures e-h, k, l drawn by  
 689 Roger Byrne. c.c. = cnemial crest; f.c. = fibular crest.

690 **a, e.** Lateral view showing worn cnemial crest at upper left.

691 **b, f.** Posterior view, showing smooth outer surface of dense bone (cf. Fig 5a, e).

692 **c, k.** Proximal end of shaft showing dense outer layer of bone and cavity filled with  
 693 coarse trabecular bone.

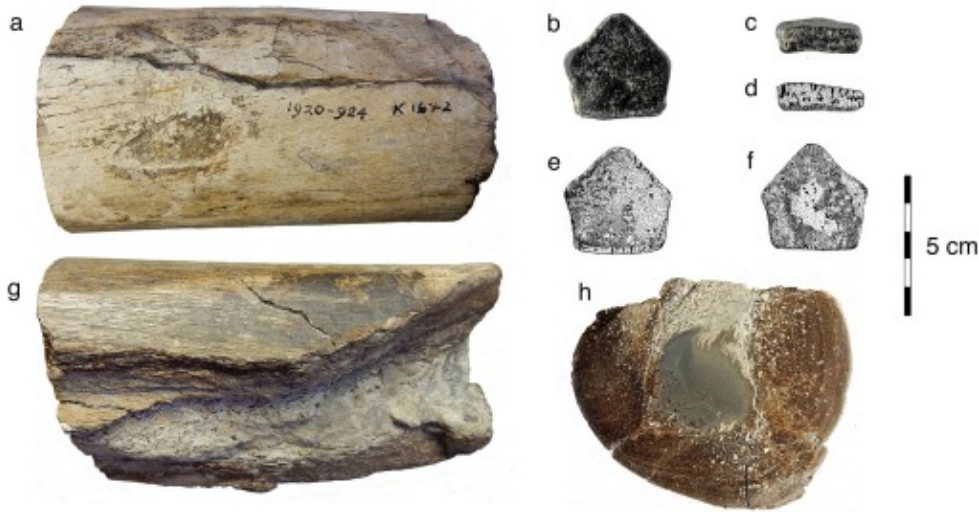
694 **d, l.** Distal end of shaft showing medullary cavity infilled with grey micritic limestone  
 695 and with spalled bone fragments in upper part.

696 **g.** Medial view.

697 **h.** Anterior view.

698 **i, j.** Left tibia of *Megalosaurus bucklandi* with section represented by K12493 outlined  
 699 (inverted figure of right tibia from Owen 1849-84, Pl. 31).

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701

702 **Fig. 5.** Objects misinterpreted as of dinosaurian origin, from Glenarm and The Gobbins.

703 Figures d-f drawn by Roger Byrne.

704 **a.** 'Outer' lateral view of BELUM K1642, possibly a fragment of ichthyosaur surangular  
705 bone, showing striated bone surface.

706 **b-f.** Polygonal 'scelidosaur osteoderm' from The Gobbins (BELUM K2015.1.54),  
707 reinterpreted as a fragment of basalt.

708 **g.** 'Inner' lateral view of BELUM K1642 showing striated outer surface and coarse  
709 trabecular bone in interior.

710 **h.** Sectioned end of BELUM K1642 showing outer layer of bone with more conspicuous  
711 porosity and LAGs than is typical of dinosaur bones (cf. K3998 and K12493, figs 3c and  
712 4c), partial fill of cavity with coarse trabecular bone, and with circumference interrupted  
713 by anomalous zone of coarse trabecular bone. Enlarged x1.5 relative to scale.